

aerosols and climate : uncertainties and the need for standardization

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From the letter of invitation:

Set the stage for Workshop that aims at developing a cross academic, industry, public stakeholder and government outlook on key aerosol metrology issues critical to furthering our understanding of aerosols in the climate *theatre*



- We are more aware of the importance of a range of aerosol characteristics and processes in the climate system.
- Quantifying the climate impact of changing the concention of individual chemical components of the atmospheric aerosol is still very uncertain.
- Quanitying the climate impact of indivicual human activities , incl. policy measures, is even more uncertain (emissions & their chemical fingerprint are uncertain), but the sign is usually known.
- Deal with uncertainty by adopting a multi-pollutant /multieffect approach (not just aerosol emissions, not just climate impacts)



- Standardization doesn't reduce uncertainty of measurements or assessements but improves their comparability.

 Policy making and policy implementation *is* about comparing (e.g. comparing climate effects with and without a policy, comparing PM values with target /limit values)

Polices often affect markets, or are themselves market based: standarization contributes to creating a level playing field.
(BC emissions tradeble like CO2 emissions?)



ATMOSPHERIC AEROSOL PARTICLES

or: Particulate Matter

size, chemical composition, morphology



EUROPEAN COMMISSION DIRECTORATE-GENERAL Joint Research Centre

chemical composition of PM2.5, worldwide (2000 – 2010)







Figure 1. TEM images of aerosol particles collected in the polluted MBL of the North Atlantic Ocean, near the Azores Islands. (a) Ammonium sulfate particles containing soot (marked by the small arrows) and fly-ash spheres (marked by the bold arrow in the lower-right corner). (b) A typical chain-like soot aggregate; the arrows point to a carbon film that connects individual spherules within the aggregate. (c) Fly-ash spheres, consisting of amorphous silica. (d) High-resolution image showing disordered graphitic layers within a soot aggregate. The background of the image is a support film of amorphous carbon onto which the particles were collected. The inset in the upper-left corner is a copy of the boxed area; the graphitic layers are highlighted in the inset.

Figure 3. Aerosol particles collected above the Southern Ocean, south of Tasmania. (a) Ammonium sulfate particles with soot inclusions (marked by arrows). The inset displays two different zone-axis SAED patterns that were obtained from a similar particle; spacings are consistent with the structure of ammonium sulfate and are indexed accordingly. (b) A typical, individual soot particle that is not enclosed by sulfate. (c) Residues and soot inclusions (arrowed) that remained after three ammonium sulfate particles were intentionally sublimed with the electron beam. (d) An ammonium sulfate particle with a small soot inclusion (arrowed). Bubbles that resulted from electron-beam damage are visible in the particles in Figures 3a and 3d. The high-resolution image shows the onion-like structure of spherules within the arrowed soot inclusion.

Posfai et al. JGR 1999



TOA Radiative Forcing for components

IPCC 4AR, 2007

Radiative Forcing Terms Climate efficacy Spatial scale (see CO, 1.0 Global High caption) Long-lived N₂O I greenhouse gases ~ 10 -1.0 - 1.2Global High 100 yrs Halocarbons Weeks to Continental Stratospheric Tropospheric Med Ozone 0.5 - 2.0100 yrs to global (-0.05) Anthropogenic Stratospheric water ~1.0 10 years Global Low vapour from CH₄ 10 -Local to Med Black carbon Surface albedo -----Land use on snow 100 yrs continental - Low Continental Med Direct effect 0.7 - 1.1 Days to global Low Total Aerosol Cloud albedo Hours Continental 1.0 - 2.0 Low effect Days to global Contrail cirrus ~ 0.6 Continental (0.01)Hours Low Natural 10 -Solar irradiance 0.7 - 1.0 Global Low 100 yrs 0 -2 2 Timescale -1 1 Scientific understanding Radiative Forcing (W m⁻²) Panel A. cooling warming 2 Long-lived greenhouse **Relative probability** gases and ozone radiative forcings Total aerosol Total anthropogenic radiative forcing radiative forcing -2 2 -3 -1 0 1 З 1 Radiative Forcing (W m-2)

Radiative forcing of climate between 1750 and 2005

Panel B.





 $\delta_{pre} \approx \delta_{nat}$

Haywood and Shine, GRL, 1995



Aerosol Optical Depth: MODIS vs. AERONET



AOD bin (550nm)



TOA Direct Radiative Forcing by Aerosols

Chung et al., JGR, 2005



Annual average: - 0,35 Wm-2 (range: - 0,6 to - 0,1 Wm-2)

Uncertainty due to

- absorption properties of the aerosol (single scattering albedo)
- vertical distribution of the aerosol





closure experiments

$$\iint Q_{sca}(\lambda, D_p, m) \frac{\pi D_p^2}{4} n(D_p, m) dm dD_p = \sigma_{sp}(\lambda)$$



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- The GAW World Calibration Centre for Aerosol Physics http://gaw.tropos.de/WCCAP/
 - Intercomparisons

- Condensation Particle Counters
 Dif. & Scan. Mobility Particle Sizers
 Nephelometers
 Absorption Spectrometers
 Hygroscopicity-TDMA
- Eu+NAm Eu+NAm+Aus Eu+NAm+Aus+Asia











Emissions of Black Carbon Aerosols



Figure 2.5 Historical BC emission trends from fossil-fuel (left) and biofuel (right) combustion from Novakov et al. (2003), Ito and Penner (2005) and Bond et al. (2007)

BC measured as

- absorbing aerosol
- elemental carbon
- fraction of PM







TOA Direct Radiative Forcing by Aerosols

Aerosol Direct Radiative Forcing





TOA 1st Indirect Radiative Forcing by Aerosols

Radiative Forcing from Cloud Albedo Effect







Integrated Assessment of Black Carbon and Tropospheric Ozone

Summary for Decision Makers





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TOA Radiative Forcing for emissions



UNEP, 2011



Global Impacts of Additional Emissions Controls on Methane and Products of Incomplete Combustion 1: CH_4 measures, 2: CH_4 +BC Group 1 measures, 3: CH_4 + all BC measures



Climate given at 2050, air quality benefits for 2030 and beyond UNEP 2011



Global Impacts of Additional Emissions Controls on Methane and Products of Incomplete Combustion 1: CH_a measures, 2: CH_a+BC Group 1 measures, 3: CH_a+ all BC measures





Climate given at 2050, air quality benefits for 2030 and beyond UNEP 2011



Standardization activities in EU by CEN

Published standards

- EN 12341: PM10 filter sampling & weighing (required by2008 directive
- EN 14907: PM2.5 filter sampling & weighing (required by2008 directive
- EN 14902: Cd, As Ni, Pb in PM
- EN 15549: BaP in PM
- prEN 15980: PAHs in PM
- CEN TC 264 currently active WGs
- WG 15: PM2.5 and PM10 (filter sampling & weighing)
- WG 21: measurement of PAHs
- WG 32: Particle number concentration
- WG 35: EC/OC in PM





TC amount measured in NIST RM 8785 by 13 labs in Europe



Fig. 1: TC amounts determined by the EUSAAR partners and reference values. Also shown are the uncertainties of these values.

- Deviations to reference cannot be explained by measurement inaccuracy (based on previous inter-comparisons)
- Large filter inhomogeneity
- No certified values for OC and EC (thermal protocol dependent)



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Aerosol & Climate science has become relevant! - Polices often affect markets, or are themselves market based: standarization contributes to creating a level playing field.

thanks



TOA Direct Radiative Forcing by Aerosols

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AeroCom (Schultz et al., 2006)

compared instantaneous direct radiative forcing due to anthropogenic aerosols produced by 9 different global models (from Europe, Japan and the USA) with detailed aerosol modules, USING THE SAME EMISSIONS

Differences in computed direct aerosol forcing (range +0.04 to -0.41Wm⁻²) are due to:

- diversity in simulated aerosol residence times
- mass extinction coefficients
- forcing efficiencies (forcing per unit optical depth)

- in clear-sky conditions: aerosol absorption, size, and surface albedo
- in cloud-sky conditions: cloud fields and relative altitudes of absorbing aerosol and clouds

⁺







Charlson et al., Science, 1992











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Radiative Forcing: bottom-up

$$RF = (E_{in} - E_{out}) - (E_{in}^{pre} - E_{out}^{pre})$$

$$RF = (E_{in} - E_{in}^{pre}) - (E_{out} - E_{out}^{pre})$$

$$RF = \Delta E_{in} - \Delta E_{out}$$

$$RF = \Delta E_{in,solar} - \Delta E_{out,GHG} - \Delta E_{out,aerosols} - \Delta E_{out,surface_albedo}$$







ATMOSPHERIC OZONE

O3 Ozone Molecule





scenarios \rightarrow emissions \rightarrow concentration fields \rightarrow climate forcing

- Model validation (against measured aerosol constituents' concentrations too)
- Model inter-comparisons

AeroCom (Schultz et al., 2006) compared instantaneous direct radiative forcing due to anthropogenic aerosols produced by 9 different global models (from Europe, Japan and the USA) with detailed aerosol modules.

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Fig. 1. The global annual mean Earth's energy budget for the Mar 2000 to May 2004 period (W m⁻²). The broad arrows indicate the schematic flow of energy in proportion to their importance.

Global Energy Flows W m⁻²



Computing current aerosol **direct** radiative forcing:

- E.g. Haywood and Shine's approximation (1995) for a given WL:



In Charlsons work, optical depth is that of anthro If you use measurements you have natural + ant

- Radiative transfer modules (e.g. Modtran) need:

- extinction vertical profile
- extinction, absorption, and asymmetry factor WL dependence

• **Computing current aerosol indirect** radiative forcing:

- mainly (?) remote sensing
- CCN number \leftrightarrow CN number size distribution, hygroscopicity



Measurements of aerosol characteristics relevant to climate forcing

are currently performed for research (rather than regulation) purposes

 \Rightarrow they are done with conscience and awareness Different flow rates in Nephelometer give same scattering measurements Different ΔT in CPCs give same number of particles with Dp>50nm

go back to a basic physical principle

 \Rightarrow they aim at getting true (and not only "comparable") values No need for standardisation if several methods can measure the same truth

E.g. - aerosol absorption from photoacoustic spectrometer =

- extinction cell (extinction - scattering)

In contrast with PM mass concentration (gravimetric methods) which give wrong data because of sampling artifacts and analytical artifacts (aerosol water uptake at 45-50% RH), and then errors have to be the same everywhere.

 \Rightarrow continuous closure checks used as strigent data quality check

Still a lot has been done regarding standardisation in the field of aerosol metrology for climate over the past decade (GAW SAG, GAW-WCCAP, ..., EMEP) Very often payed by European Commission





- The GAW – World Calibration Centre for Aerosol Physics http://gaw.tropos.de/WCCAP/

- Intercomparisons



93

Audits (on-site check of sampling configuration, instruments, etc...)

2003

2008

2003

2004

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2010

2010?

Tamanrasset/Assekrem Algeria Hohenpeissenberg Germany Anmyeon South-Korea Pallas Finland Danum Valley Malaysia Mace Head Ireland Waliguan-Beijing China Danum Valley Malaysia Danum Valley Malaysia Cape Point South Africa Ny-Alesund Norway Beo-Moussala Bulgaria Tenerife, Spain Izana Point Barrow Alaska, USA Monte Cimone Italy Harwell UK Mauna Loa Hawaii, USA Finokalia Crete, Greece Cabauw The Netherlands Bukit Koto Tabang Indonesia Vavihill Sweden Cape Grim Australia Shangdianzi China Danum Valley Malaysia Hyytiälä Finland Birkenes Norway Finokalia Crete, Greece Preila Lithuania Aspvreten Sweden Melpitz Germany Ispra Italy JFJ, PdD, Kosetice, K-Puszta

I can make a GANTT-like diagram with this too, highlighting stations from the 5 continents



scenarios \rightarrow emissions \rightarrow concentration fields \rightarrow climate forcing

- Model validation (against measured aerosol constituents' concentrations too)

- inorganics: nitrate, sulfate, (ammonium), etc...

- ∃ reference materials (NIST **S**RM 2694, IRMM xxx, ...)

 ∃ yearly or twice yearly regular inter-comparisons GAW - QA/SAC-Americas: 78 labs worldwide (<u>http://qasac-americas.org/</u>) EMEP- CCC: 51 labs worldwide (<u>http://tarantula.nilu.no/projects/ccc/</u>)

- carbonaceous aerosol

- reference material (NIST RM 8785)?

- yearly inter-comparison with EUSSAR project (18 Eu labs)

- punctual inter-continental inter-comparison EnvCan, IMPROVE, EUSAAR





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$$\sigma_{sca}(\lambda,\bar{r},t) = \iint Q_{sca}(\lambda,D_p,m) \frac{\pi D_p^2}{4} N(D_p,m,\bar{r},t) dm dD_p$$





$$CCN(S_c, \bar{r}, t) = \iint H[D_p - D_{p,c}(S_c, \varepsilon)] N(D_p, \varepsilon, \bar{r}, t) d\varepsilon dD_p$$





The GAW – World Calibration Centre for Aerosol Physics <u>http://gaw.tropos.de/WCCAP/</u>

- Intercomparisons
- Audits
- Travelling reference instruments (for especially remote stations)





Definitions

What is a Standard?

A standard is an established norm or requirement

It is usually a formal document that establishes uniform technical criteria, methods, processes and practices

It may be developed privately or unilaterally, for example by a corporation, regulatory body, military

Simply They define how things should be done



$$\sigma_{sca}(\lambda,\bar{r},t) = \iint Q_{sca}(\lambda,D_p,m) \frac{\pi D_p^2}{4} N(D_p,m,\bar{r},t) dm dD_p$$





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BLACK CARBON PARTICLES

"Black Carbon" or "Elemental Carbon" is:

- carbon-containing particulate matter, resulting from incomplete combustion
 - it withstands high temperatures
 - it is black, hence absorbs light
- emitted together with CO₂, CO, NMHC, organic particlate matter)









Anthropogenic forcing??



Figure 2.13. Estimates of the direct aerosol RF since the TAR from observational based studies, independent modelling studies, and AEROCOM results with identical aerosol and aerosol precursor emissions. GISS_1 refers to a study employing an internal mixture of aerosol, and GISS_2 to a study employing an external mixture.



Figure 2.14. Radiative forcing due to the cloud albedo effect, in the context of liquid water clouds, from the global climate models that appear in Table 2.7. The labels next to the bars correspond to the published study; the notes of Table 2.7 explains the species abbreviations listed on the left hand side. Top panel: results correspond to models that consider a limited number of species, primarily anthropogenic sulphate (S). Bottom panel: results correspond to studies that include a variety of aerosol compositions and mixtures; the estimates here cover a larger range than those in the top panel. Chen and Penner (2005) present a sensitivity

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Koch et al., 2009



On air quality

Black carbon and ozone in the lower atmosphere are harmful air pollutants affecting health of humans and ecosystems

Black carbon, a component of particulate matter, and ozone both lead to premature deaths worldwide. (PM2.5: ~2 million per year, WHO)

Ozone is also the most important air pollutant, responsible for reducing crop yields, and thus affects food security.

On global and regional climate ?



On global and regional climate

Changes in their burdens over the 20th Century result in an expected equilibrium global warming of 0.0-0.8 C due to BC and 0.1-0.4 C due to O_3 . (The equilibrium warming expected from CO2 is about 1.3 C.)

Atmospheric heating by BC disturbs tropical rainfall and regional circulation patterns such as the Asian monsoon.

Black carbon's darkening of snow and ice surfaces increases their absorption of sunlight, which, along with atmospheric heating, exacerbates melting of snow and ice around the world, including in the Arctic, the Himalayas.



change in BC deposition: 2005 to 2030 reference scenario





UNEP, 2011

%



Avoidance of 0.2-1.8 million premature deaths per year , in North America, East Asia, SE Asia & Pacific, While in increase by 0.1-2.0 million in South, West & Central Asia and Africa..

Decreasing production of wheat, rice corn and soy by 7 to 120 million tonnes across Asia with an associated economic loss of US1 - 20 billion. Increasing crop yields in US and Europe.

The compensating warming and cooling impacts of changes in BC, ozone and other aerosols lead to a small net warming of less than 0.1 C globally

Black carbon's darkening of snow and ice and atmospheric heating, keeps exacerbating melting of snow and ice in the Himalayas



Conclusions so far:

based on our knowledge of the behaviour of BC and ozone in the atmosphere, and of their impacts on air quality, And on lobal and regional climate in teh near term

it makes a lot of sense to look for measures that specifically target the emissions of BC and ozone precursors in order to solve a range of important problems at once.





Closure experiments

scattering, absorpttion

cloud activation:

$$\sigma_{sp}(\lambda) = \iint Q_{sca}(\lambda, D_p, m) \frac{\pi D_p^2}{4} n(D_p, m) dm dD_p$$
$$CCN(S_c) = \iint H[D_p - D_{p,c}(S_c, \varepsilon)] n(D_p, \varepsilon) d\varepsilon dD_p$$
$$PM_x = \int_0^x \int \rho(D_p) \frac{\pi D_p^3}{6} n(D_p, \rho) d\rho dD_p$$

particulate matter

assuming internal mixing of chemical species into an average one

$$n(D_p,\varepsilon) = n(D_p)\delta(\varepsilon - \overline{\varepsilon})$$

with δ the Kronecker delta function

assuming external mixing of i chemical species

$$n(D_p,\varepsilon) = \sum_{i=1}^{k} n_i(D_p) \delta(\varepsilon - \varepsilon_i)$$



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Aerosol & Climate science has become relevant! - Policy making and policy implementation about camparing (e.g. climate effect with and without a policy, comparing PM

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$$RF = (E_{in} - E_{out}) - (E_{in}^{ref} - E_{out}^{ref})$$

reference = pre-industrial era

hence
$$(E_{in}^{ref} - E_{out}^{ref}) \approx 0$$

and $RF = (E_{in} - E_{out})$







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Differences in computed direct aerosol forcing (range +0.04 to -0.41Wm⁻²) are due to:

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 +
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- For any statement that I going to make there is an expert in the room that can go much mre n depth, or even challenge it
- Steve: father if not grand-father: without a beard mor like a father
- When interest of standardization organizations > relevance!
- PM/healt already relevant: legislation > lavel playing field money (even though this community knows much less aerosol,but there is epidemiological evidence!) (do we havedirect evidence of impacts of aerosols on climate:, and that it is important?)
- How can we go from microscopic compexity to values regional or glonal impacts on radiation balece and climate? Do we need to consider all that complexity, to come to usefull values?
 - Not yet legislation for aerosols and climate, scientific need (but e.g. BC in CDM is discussed ?) anyway relevance
 - Ozone vs PM, PM external internal mixing: onc you ave O3, RF easy to calculate.once you have PM dot easy to calculate
 - TOA forcing, RF = delta(present) delta (pre-industrial) > role of knowing pre-industrial values
 - Overall Earth radiation balance
 - Delta aerosols, delta GHG.delat alsbedl
 - IPCC usual graph (not policy relevant), immediately show other IPCC (more relevant) > Shindell
 - At this point usefull to make a distinction between assessement of present day effects of PM on climate (relevant to monitor compliance with legilation: standardization) predicting future impacts of PM on climate (policy development)
 - Paper Charlson et al. / Haywood & Shine
 - based on secondary parameters: problem of pre-industrial
 - Based on primar parameters related to emisioins
 - What did we learn? Include TOA > TOA+ surface+atmospheric forcing >regional impacts
 - We know a lot about aerosols, compared to what health people know or want to consider about PM. Still the latter made it to policy and hence standardization: research becomes different ...
 - Ibe exampels of standardization work in air quality work
 - Two communities are any way mergibg: e.g. UNEP assessement.
 - RF of BC emissions (IGAC, telephone conference India, US, Nairobi, Europe)
 - Effect on Climate, Health Crops. Integrated > uncertainties are less problematic.
 - Conclusion?